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OF COLUMBIA UNIVERNITY

Fallocates, New York 19964

THE INTEGRAL SOLVELION OF THE SOUND PIELS
THE A BULLTHANKER'S STOUTH-SOLID HALL ASPACE.
THE ACKERLOAD COMPONATIONS FOR LOW-PREQUENTY PROPAGATION
IN THE ANOTIC OCEAN

by

Wearn W. Rutechtle

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Territori Report No. 1

Contract 100019-67-5-0108-0016 with the Office of Navel Basescon

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Corrections

Page 7, line 3, should read as follows:

in terms of elements of

Page 13, line 4, should read as follows:

Page 16, line 8, should read as follows:

$$\begin{bmatrix} \hat{w}_{n-1} \\ (p_{22})_{n-1} \end{bmatrix} = a_{n-1} \cdot \cdot \cdot b_{n-3} a_{n-4} \cdot \cdot \cdot b_1 \begin{bmatrix} \hat{w}_0 \\ 0 \end{bmatrix}$$

Page 19, line 1, should read as follows:

$$A = A_{s_2} A_{s_1} = a_{n-1} a_{n-2} \cdots b_{n-10} \cdot a_1$$

Page 22, line 7, should read as follows:

at depth 2 at the bottom of the D₁ layer is +D₁(Γ , Z, K, t) =

Page 22, line 10, should read as follows:

$$A_{D_i} = \alpha_{D_i} \alpha_{D_i-1} \cdots \alpha_i$$

Page 28, line 2, should read as follows: but from the relation

$$AA^{1} = I$$
 $A_{11}A_{22} - A_{12}A_{21} = 1$

Page 41, second line, should read as follows:

$$\left\{
\frac{2\sqrt{2\pi^{2}}\sqrt{\omega^{2}}}{\rho_{N}c^{8}\sqrt{2}\sqrt{2}\sqrt{c^{2}A_{11}}+Im(-A_{21})}\right\} = \omega_{m}$$

$$C = C_{m}$$

LAMONT-DOHERTY GEOLOGICAL OBSERVATORY

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THE INTEGRAL SOLUTION OF THE SOUND FIELD
IN A MULTILAYERED LIQUID-SOLID HALF-SPACE
WITH NUMERICAL COMPUTATIONS FOR LOW-FREQUENCY PROPAGATION
IN THE ARCTIC OCEAN

bу

Henry W. Kutschale

CU1-1-70

Technical Report No. 1

Contract N00014-67-A-0108-0016 with the Office of Naval Research

Work done on behalf of the U.S. Naval Ordnance Laboratory, White Oak, Silver Spring, Maryland

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February 1970

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ABSTRACT

This report develops by matrix methods the integral solution of the wave equation for point sources of harmonic waves in a liquid layer of a multilayered liquid-solid half space in a form convenient for numerical computation on a high-speed digital computer. Only the case is considered of a high-speed liquid bottom underlying the stack of layers. The integral over wave number has singularities in the integrand and is conveniently transformed into the complex plane. By a proper choice of contours, complex poles are displaced to an unused sheet of the two-leaved Riemann surface, and the integral solution for the multilayered system reduces to a sum of normal modes plus the sum of two integrals, one along the real axis and the other along the imaginary axis. Both integrals are evaluated by a Gaussian quadrature formula. Sample computations are presented for low-frequency propagation in the Arctic Ocean sound channel. These are preliminary computations and the ice layer, which averages three meters in thickness, is not included in the layered system. The effects of the ice layer on propagation are currently under investigation.

FIGURES

- Figure 1. Multilayered half-space,
- Figure 2. Contours for integration.
- Figure 3. Variation of sound velocity with depth for Model A.

 Table 2 gives additional parameters for this model.
- Figure 4. Computations for Model A. Phase and group-velocity dispersion and excitation function of pressure dependent only on layering.
- Figure 5. Computations for Model A. Variation with range of the absolute value of pressure for the normal-mode contribution of the sound field. Source depth 150 m. Hydrophone depth 50 m. Source frequency 10 Hz. Source pressure amplitude 1 dyne/cm² re 1m.
- Figure 6. Computations for Model A. Variation with range of absolute value of pressure for the integral contribution of the sound field. Source and detector same as for Figure 5.
- Figure 7. Computations for Model A. Variation with range of absolute value of pressure of the total sound field. Source and detector same as for Figure 5.
- Figure 8. Same as Figure 5 but computations carried to longer range.
- Figure 9. Same as Figure 6 but computations carried to longer range.
- Figure 10. Same as Figure 7 but computations carried to longer range.

DEFINITION OF SYMBOLS

Om; compressional-wave velocity in the m-th layer Bm; shear-wave velocity in the m-th layer hm; thickness of the m-th layer ₹ ; vertical coordinate ; range between source and detector t; time (W); angular frequency C; phase velocity k; wave number i; V-T \sum_{m} ; imaginary part of a complex number x + iy $\phi_{\mathbf{m}}$; velocity potential in the m-th layer ; normal stress in the m-th layer parallel to z axis ; pressure in the m-th liquid layer horizontal particle displacement in the m-th layer Wm; vertical particle displacement in the m-th layer horizontal particle velocity in the m-th layer wm; vertical particle velocity in the m-th layer Cm; density in the m-th layer Ps; density at the source

Kam= W/Am

Kpm W/Bm

 $\int_{\mathbf{n}}$; Bessel function of order 0 \bigvee_{0} ; Y bessel function of order 0 K; K Bessel function of order 0 Hankel function of the first kind of order 0 Hankel function of the second kind of order 0 Bam = VK2-K2 , KKKam Bam =-iVk2-k21 , k> kam BBm = VKBm- K2, K < KBM BBm = - i V K2 - KBm > K > KBm Pm = hm Bam $Q_m = h_m \beta_{\beta m}$ $V_{\alpha_m} = \sqrt{\frac{c}{\alpha_m}^2 - 1}$, $c > \alpha_m$ ram = - (/ 1 - (c / 2) , c < am rpm=1(C) -1 , C>Bm rpm = - (1 - (CBM) , C (BM

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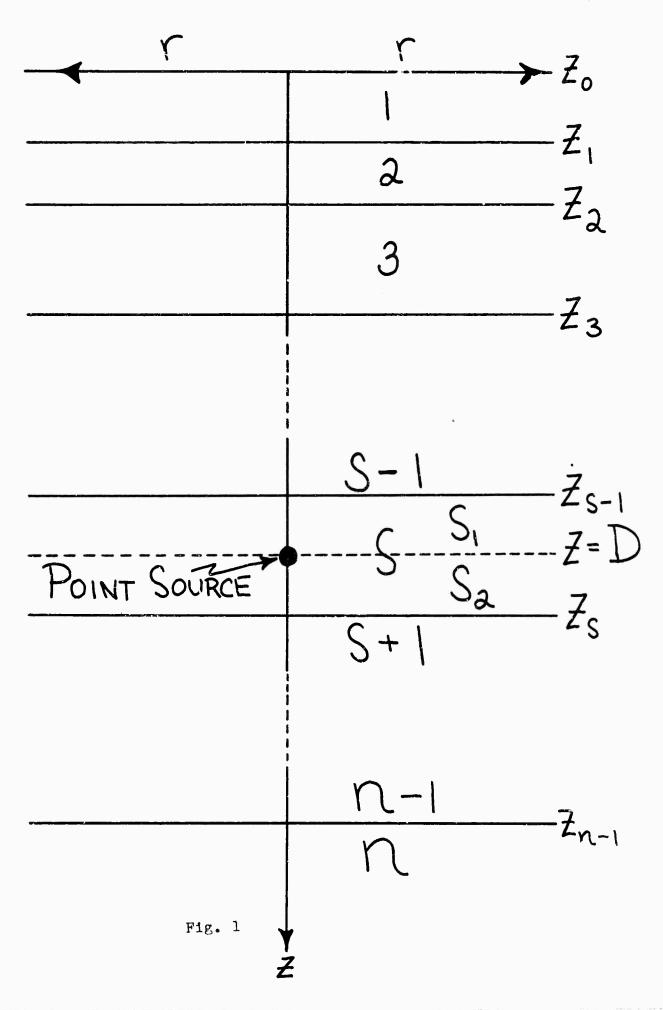
INTRODUCTION

This report develops the integral solution of the wave equation for point sources of harmonic waves in a liquid layer of a multilayered liquid-solid half space in a form convenient for numerical computation on a high-speed digital computer. Only the case of a liquid half-space underlying the stack of layers is considered. Furthermore, it is assumed that the speed of sound in this last layer of infinite thickness is greater than the speed of sound in liquid layers immediately above it. The problem of immediate concern is low-frequency propagation from harmonic sources in the central Arctic Ocean at ranges up to ten times the water depth in shallow and intermediate water depths. Solutions are developed for the pressure perturbations detected by a hydrophone at depth or the ice vibrations detected by a seismometer on the ice surface. A modification of the formulation is possible to include harmonic vibration sources on or in the ice, although it appears at the present time that surface sources are of limited application (Greene, 1968).

The solution of the wave equation presented here, based on the Thomson-Haskell matrix method (Thomson, 1950; Haskell, 1953), follows Harkrider (1964) for the solution of the wave equation in an n-layered solid half space. Layer matrices of the type given by Dorman (1962) for computing dispersion in an n-layered liquid - solid half-space are used for the liquid layers. An application of the theorem that the inverse of the

product matrix for the layered system above the source has the same form as the inverse of a layer matrix reduces the integrand of the integral solution to a simple form in terms of product matrices derived in the source-free, plane-wave case. integral over wave number from zero to infinity has singularities in the integrand and is, therefore, conveniently transformed into the complex $\zeta = k + i\tau$ plane. Performing the contour integration in the z plane and noting that branch line integrals corresponding to branch points generated by each layer matrix except the last are zero, we are left with an expression analogous to that of Sorensen (1959) and Leslie and Sorensen (1961) for the two-layer liquid half-space with a high-speed bottom. By the proper choice of contours, complex poles are displaced to an unused sheet of the two-leaved Riemann surface, and the integral solution for the multilayered system reduces to a summation of normal modes plus the sum of two integrals over wave number, one along the real z axis and the other along the imaginary ; axis. Both of these integrals are conveniently computed by a Gaussian quadrature formula.

The physical interpretation of the final solution is straightforward. The normal-mode terms correspond to waves trapped in the Arctic sound channel; that is, refracted and surface reflected (RSR) sounds and reflected sounds incident on the bottom beyond the critical angle. The integral over wave number along the real axis corresponds to waves incident



on the bottom at angles greater than the critical angle of total reflection. Since energy leaves the guide continuously as waves travel down the guide, this term is of importance only to ranges of about ten times the water depth. The contribution to the sound field of the integral along the imaginary axis is of importance only very near the source since the integrand decays exponentially with range and wave number.

The usefulness of the normal mode terms for describing long-range explosive sound transmission in the central Arctic Ocean is shown by Kutschale (1969). In that work the present formulas for the normal modes were extended to explosive sources and the Fourier integral for each mode was evaluated by the principle of stationary phase (Pekeris, 1948). Attenuation by the rough ice boundaries was also included, although omitted here.

FORMAL SOLUTION

Source Free Case

Consider the n-layered interbedded liquid-solid half-space shown in Figure 1 in which the last layer of infinite thickness is liquid and has a higher sound velocity than liquid layers immediately above it; that is, a high-speed bottom. A point source of harmonic waves is located in one of the liquid layers. The velocity potential in the m-th layer, which is

assumed to be liquid, satisfies the wave equation

$$\nabla^2 \phi_m - \frac{1}{d^2} \phi_m = 0.$$

In this layer

$$(\phi_{22})_{m} = -\phi_{m} = \rho_{m} \frac{\partial \phi_{m}}{\partial t}, \quad \dot{u}_{m} = \frac{\partial \phi_{m}}{\partial z}, \quad \dot{u}_{m} = \frac{\partial \phi_{m}}{\partial r}.$$

Solutions for pm, wm and (b) are given by

$$\dot{w}_{m}(r,z,t,k) = (\hat{w}_{m}(z)) J(kr)e^{i\omega t} dk$$

$$(\dot{p}_{zz})_{m}(r,z,t,k) = (\hat{p}_{zz})_{m}(z) J(kr)e^{i\omega t} dk$$

Separation of variables yields

$$\widehat{W}_{m}(z)e^{i\omega t} = \frac{\partial \widehat{\rho}_{m}e^{i\omega t}}{\partial z} = \left(i\sqrt{k_{a_{m}}^{2}-k^{2}}A_{m}e^{i\sqrt{k_{a_{m}}^{2}-k^{2}}z}\right)e^{i\omega t}$$

$$-i\sqrt{k_{a_{m}}^{2}-k^{2}}A_{m}e^{-i\sqrt{k_{a_{m}}^{2}-k^{2}}z}\right)e^{i\omega t}$$

where A_{m} and A_{m}^{\bullet} are constants.

Placing the origin at the (m-1)-st interface we get

$$\widehat{w}_{m-1} = i\sqrt{k_{d_m}^2 - k^2} \left(A_m - A_m \right)$$
(1)

$$(\hat{p}_{zz})_{m-1} = i \omega p_m (A_m + A_m)$$
 for $z = 0$

and

$$\hat{w}_{m} = -i \sqrt{k_{dm}^2 - k^2} (A_m + A_m) \sin P_m + i \sqrt{k_{dm}^2 - k^2} (A_m - A_m) \cos P_m$$
 (2)

for
$$Z = h_{m}$$

Substituting expressions for $[A_m - A_m^*]$ and $[A_m + A_m^*]$ from equations (1) in equations (2) yields

equations (1) in equations (2) yields
$$\widehat{W}_{m} = \widehat{W}_{m-1} COSP_{m} - (\widehat{p}_{zz})_{m-1} \frac{\sqrt{k_{x_{m}}^{2} - k^{2}}}{i \omega p_{m}} SINP_{m}$$

$$(\widehat{p}_{zz})_{m} = -\widehat{W}_{m-1} \frac{\omega p_{m}}{i \sqrt{k_{x_{m}}^{2} - k^{2}}} SINP_{m} + (\widehat{p}_{zz})_{m-1} COSP_{m}$$

or in matrix notation $\begin{bmatrix}
\hat{W}_{m} \\
\hat{P}_{zz} \\
m
\end{bmatrix} = \begin{bmatrix}
\cos P_{m} \\
\frac{i \sqrt{k_{a_{m}}^{2} - k^{2}}}{\sqrt{k_{a_{m}}^{2} - k^{2}}} & \cos P_{m} \\
\frac{i \sqrt{k_{a_{m}}^{2} - k^{2}}}{\sqrt{k_{a_{m}}^{2} - k^{2}}} & \cos P_{m}
\end{bmatrix}$

$$\left[\begin{array}{c} \hat{\omega}_{m-1} \\ (\hat{p}_{zz})_{m-1} \end{array}\right] =$$

$$\begin{bmatrix}
 (a_{m})_{11} & (a_{m})_{12} \\
 (a_{m})_{21} & (a_{m})_{22}
 \end{bmatrix}
 \begin{bmatrix}
 (\hat{p}_{22})_{m-1} \\
 (\hat{p}_{22})_{m-1}
 \end{bmatrix}$$

In general for an n-layered liquid half-space

$$\begin{bmatrix} \hat{w}_{n-1} \\ (+_{zz})_{n-1} \end{bmatrix} = \begin{bmatrix} A_{11} & A_{12} \\ A_{21} & A_{22} \end{bmatrix} \begin{bmatrix} \hat{w}_{0} \\ (+_{zz})_{0} \end{bmatrix}$$
or
$$\begin{bmatrix} \hat{w}_{n-1} = A_{11}\hat{w}_{0} + A_{12}(+_{zz})_{0} \\ (+_{zz})_{0} \end{bmatrix}$$
where
$$\begin{bmatrix} \hat{\rho}_{n-1} \hat{\rho}_{zz} = A_{21}\hat{w}_{0} + A_{22}(+_{zz})_{0} \\ A_{21} & A_{22} \end{bmatrix} = \alpha_{n-1}\alpha_{n-2} \cdots \alpha_{1}$$

$$= \begin{bmatrix} A_{11} & A_{12} \\ A_{21} & A_{22} \end{bmatrix} = \alpha_{n-1}\alpha_{n-2} \cdots \alpha_{1}$$

$$= \begin{bmatrix} (\alpha_{n-1})_{11} & (\alpha_{n-1})_{12} \\ (\alpha_{n-1})_{21} & (\alpha_{n-1})_{22} \end{bmatrix} \cdots \begin{bmatrix} (\alpha_{1})_{11} & (\alpha_{1})_{12} \\ (\alpha_{1})_{21} & (\alpha_{1})_{22} \end{bmatrix}$$

since the vertical particle velocity and pressure are continuous across each interface. For the n-th layer

$$\Delta_{n} - \Delta_{n}' = \frac{\hat{w}_{n-1}}{i\sqrt{k_{d_{n}}^{2} - k^{2}}} = \frac{A_{11}\hat{w}_{o}}{i\sqrt{k_{d_{n}}^{2} - k^{2}}} + \frac{A_{12}(\hat{p}_{zz})_{o}}{i\sqrt{k_{d_{n}}^{2} - k^{2}}}$$

$$A_{n} + A_{n} = \frac{(\hat{\rho}_{zz})_{n-1}}{i\omega\rho_{n}} = \frac{A_{21}\hat{\omega}_{0}}{i\omega\rho_{n}}$$
or in matrix notation
$$A_{22}(\hat{\rho}_{zz})_{0}$$

$$A_{n} - A_{n}$$

$$A_{n} - A_{n}$$

$$A_{n} + A_{n} = E^{-1}(\hat{\rho}_{zz})_{n-1}$$

$$E^{-1}(\hat{\rho}_{zz})_{n-1}$$

$$E^{-1}(\hat{\rho}_{zz})_{n-1}$$

Applying boundary conditions that the pressure vanishes at the surface and that no upgoing waves travel from infinity, $(p_{ZZ})_0 = 0$, $A_n = 0$, the relation (3) reduces to

$$\begin{bmatrix} -A_n' \\ A_n' \end{bmatrix} = E^- A \begin{bmatrix} \hat{w}_0 \\ 0 \end{bmatrix}$$

Consider now a solid layer between two liquid layers or at the surface of the laminated halfspace. For this layer we may write (Thomson, 1950; Dorman, 1962)

where the fourth element of the column vector is the tangential stress which is zero at a solid-liquid boundary. The matrix elements are given in the appendix and may be derived following Haskell (1953). Equation (4) yields the three equations.

$$\hat{u}_{m-1} = -\frac{(\alpha_m)_{42}}{(\alpha_m)_{41}} \hat{w}_{m-1} - \frac{(\alpha_m)_{43}}{(\alpha_m)_{41}} (\hat{p}_{zz})_{m-1}^{(5)}$$

$$\hat{w}_{m} = (Q_{m})_{21}\hat{Q}_{m-1} + (Q_{m})_{22}\hat{w}_{m-1} + (Q_{m})_{23}\hat{f}_{23}^{(6)}$$

Substituting $\widehat{\mathbf{Q}}_{m-1}$ from equation (5) in equations (6) and (7) yields

$$\dot{w}_{m} = \left[(a_{m})_{22} - \frac{(a_{m})_{21}(a_{m})_{42}}{(a_{m})_{41}} \right] \dot{w}_{m-1} +$$

$$\left[\frac{(a_m)_{23} - \frac{(a_m)_{21}(a_m)_{43}}{(a_m)_{41}}\right](\hat{f}_{z\bar{z}})_{m-1}$$

which is the same as the matrix relation for a liquid layer.

In general, then, for interbedded solids and liquids

$$\begin{bmatrix} \hat{w}_{n-1} \\ \hat{p}_{22} \\ n-1 \end{bmatrix} = a_{n-1} \cdots b_{n-3} a_{n-4} \cdots b_{0} \begin{bmatrix} \hat{w}_{0} \\ 0 \end{bmatrix}$$

$$= A \begin{bmatrix} \hat{w}_{0} \\ 0 \end{bmatrix}.$$

Point Source of Harmonic Waves in a Liquid Layer.

Divide the source layer into two layers as shown in Figure 1. At z=D the pressure is continuous. The vertical particle velocity is continuous everywhere in the plane defined by z=D except at the point source where the liquid above and below the source moves in opposite directions. This may be expressed by writing

$$\delta(\hat{w})_s = 2K$$
.

In matrix notation for the source layer

$$\begin{bmatrix} \hat{w}_{S_{2}}(D) \\ (\hat{\varphi}_{zz})_{S_{2}}(D) \end{bmatrix} = \begin{bmatrix} \hat{w}_{S_{1}}(D) \\ (\hat{\varphi}_{zz})_{S_{2}}(D) \end{bmatrix} + \begin{bmatrix} \delta(\hat{w})_{S} \\ 0 \end{bmatrix}.$$
(8)

For the layers below z = D we have

$$\begin{bmatrix}
\hat{w}_{n-1} \\
\hat{\varphi}_{zz} \\
n-1
\end{bmatrix} = A_{s_2} \begin{bmatrix}
\hat{w}_{s_2}(D) \\
\hat{\varphi}_{zz} \\
s_2
\end{bmatrix}$$
(9)

and for the layers above z=D

$$\begin{bmatrix} \hat{w}_{s,(D)} \\ \hat{p}_{t} \end{bmatrix} = A_{s,(D)}$$

$$\begin{bmatrix} \hat{w}_{o} \\ \hat{p}_{t} \end{bmatrix}$$

where $A_{s_2} = a_{n-1} a_{n-2} \cdots a_{s_2}$ and $A_{s_1} = a_{s_1} a_{s_1-1} \cdots a_{s_n}$.

Returning now to the relation for the n-th liquid layer

we write from equations. (8), (9), and (10)
$$\begin{bmatrix}
-A_n \\
A_n'
\end{bmatrix} = E^{-1} \begin{bmatrix} \widehat{W}_{n-1} \\ (\widehat{p}_{22})_{n-1} \end{bmatrix} = E^{-1}A_{S_2} \begin{bmatrix} \widehat{W}_{S_2}(D) \\ (\widehat{p}_{22})_{S_2}(D) \end{bmatrix} + \begin{bmatrix} \delta(\widehat{W})_S \end{bmatrix}$$

$$= E^{-1}A_{S_2} \begin{bmatrix} \widehat{W}_{S_1}(D) \\ (\widehat{p}_{22})_{S_1}(D) \end{bmatrix} + \begin{bmatrix} \delta(\widehat{W})_S \end{bmatrix}$$

$$= E^{-1}A_{S_2} \begin{bmatrix} A_{S_1} \begin{bmatrix} \widehat{W}_{S_1} \\ S_1 \end{bmatrix} \\ A_{S_2} \begin{bmatrix} \widehat{W}_{S_1}(D) \\ S_1 \end{bmatrix} + \begin{bmatrix} \delta(\widehat{W})_S \end{bmatrix}$$

$$= E^{-1} \begin{bmatrix} A_{S_2} A_{S_1} \begin{bmatrix} \widehat{W}_{S_1} \\ S_1 \end{bmatrix} + A_{S_2} \begin{bmatrix} \delta(\widehat{W})_S \end{bmatrix}$$

$$= E^{-1} \begin{bmatrix} A_{S_2} A_{S_1} \begin{bmatrix} \widehat{W}_{S_1} \\ S_1 \end{bmatrix} + A_{S_2} \begin{bmatrix} \delta(\widehat{W})_S \end{bmatrix}$$
(11)

In terms of the inverse matrix $A_{s_1}^{-1}$ of A_{s_1} (11) may be written

$$\begin{bmatrix} -A_{n} \\ A_{n} \end{bmatrix} = E^{-1}A \left\{ \begin{bmatrix} \hat{w}_{0} \\ 0 \end{bmatrix} + A_{s_{1}}^{-1} \begin{bmatrix} \delta(\hat{w})_{s} \\ 0 \end{bmatrix} \right\}_{(12)}$$

where
$$A = A_{s_a}A_{s_1} = a_{n-1}a_{n-2}\cdots b_{n-10}\cdots a_0$$

$$\begin{bmatrix} X \\ Y \end{bmatrix} = \begin{bmatrix} \widehat{w}_0 \\ 0 \end{bmatrix} + A_{S_i}^{-1} \begin{bmatrix} \delta(\widehat{w})_s \\ 0 \end{bmatrix}.$$

Harkrider (1964) has shown that for layered solids the inverse of the product matrix has the same form as the inverse of the layer matrices. Employing an extension of this theorem to interbedded liquids and solids we get

$$A_{s_i}^{-1} = \begin{bmatrix} (A_{s_i})_{22} & -(A_{s_i})_{12} \\ -(A_{s_i})_{21} & (A_{s_i})_{11} \end{bmatrix}$$

Therefore
$$\begin{bmatrix}
X \\
Y
\end{bmatrix} = \begin{bmatrix}
\hat{W}_{0} \\
0
\end{bmatrix} + \begin{bmatrix}
(A_{s_{i}})_{22} & -(A_{s_{i}})_{12} \\
-(A_{s_{i}})_{21} & (A_{s_{i}})_{11}
\end{bmatrix}
\begin{bmatrix}
\delta(\hat{W})_{s} \\
0
\end{bmatrix}
= \begin{bmatrix}
\hat{W}_{0} \\
0
\end{bmatrix} + \begin{bmatrix}
(A_{s_{i}})_{22} & \delta(\hat{W})_{s} \\
-(A_{s_{i}})_{21} & \delta(\hat{W})_{s}
\end{bmatrix}$$
or

$$X = \hat{w}_{o} + (A_{s_{i}})_{22} \delta(\hat{w})_{s}$$

$$Y = -(A_{s_{i}})_{21} \delta(\hat{w})_{s}.$$
(13)

We now define
$$J = E^{-1}A = \begin{bmatrix} J_{11} & J_{12} \\ J_{21} & J_{22} \end{bmatrix}$$
.

$$\begin{bmatrix} -A_{1}' \\ -A_{2}' \end{bmatrix} = J\begin{bmatrix} X \\ X \end{bmatrix} = \begin{bmatrix} J_{11}X + J_{12}Y \\ J_{21}X + J_{22}Y \end{bmatrix}$$
 or

$$-A'_{n} = J_{11}X + J_{12}Y$$

$$A'_{n} = J_{21}X + J_{22}Y.$$
(14)

Adding equations (14) we get

$$0 = (J_{11} + J_{21}) X + (J_{12} + J_{22}) Y$$
 and if we

let $T = J_{11} + J_{21}$ and $V = J_{12} + J_{22}$ and solve for X we have

$$X = -\frac{VY}{T}.$$
(15)

Carrying out the matrix multiplication $J = E^{-1} A$, expressions for T and V are given by

$$T = \frac{A_{11}}{i\sqrt{k_{a_n}^2 - k_{a_1}^2}} + \frac{A_{a_1}}{i\omega\rho_n}$$

$$V = \frac{A_{1a}}{i\sqrt{k_{a_n}^2 - k_{a_1}^2}} + \frac{A_{2a}}{i\omega\rho_n}.$$
(16)

From the first equation of (13)

from equation (15)

$$\hat{w}_{o} = -\frac{VY}{T} - (A_{S_{1}})_{22} \delta(\hat{w})_{S}$$

Hence, from the second equation of (13) and equations (16)

$$\hat{W}_{0} = \left(\frac{A_{12}}{i\sqrt{K_{u_{n}}^{2} - K_{u}^{2}}} + \frac{A_{22}}{i\omega\rho_{n}}\right)(A_{S_{1}})_{2},\delta(\hat{w})_{S}$$

$$\frac{A_{11}}{i\sqrt{K_{u_{n}}^{2} - K_{u}^{2}}} + \frac{A_{21}}{i\omega\rho_{n}}$$

$$-(A_{s_1})_{22}\delta(\hat{w})_{s}$$

Therefore the integral solution for the surface verticle particle velocity is $\dot{w}_0(\zeta, z, k, t) =$

$$\left(\frac{A_{12}}{cvt} + \frac{A_{22}}{cvp_{n}}(A_{s_{1}})_{21} - \frac{A_{11}}{cvk_{n}^{2} - k^{21}} + \frac{A_{21}}{cwp_{n}}(A_{s_{1}})_{22}}{A_{11}} + \frac{A_{21}}{cwp_{n}}(A_{s_{1}})_{22} + \frac{A_{21}}{cwp_{n}}(A_{s_{1}})_{22}}\right)$$

$$2 J_{0}(kr)kdk$$

and from the formula

$$(\hat{P}_{22})_{i} = (A_{D_{i}})_{2} \hat{W}_{0} = -\hat{P}_{D_{i}} \text{ the integral solution for pressure}$$
at depth D_{i} is $\hat{P}_{D_{i}}(\Gamma, Z, k, t) =$

$$- \left(\frac{e^{i\omega t}}{e^{i\omega t}} \right) \left(\frac{A_{12}}{i\sqrt{k_{\alpha_{in}}^{2} - k^{2}}} + \frac{A_{22}}{i\omega p_{in}} \right) (A_{S_{i}})_{2i}$$

$$\frac{A_{1i}}{i\sqrt{k_{\alpha_{in}}^{2} - k^{2}}} + \frac{A_{2i}}{i\omega p_{in}} (A_{S_{i}})_{22}$$

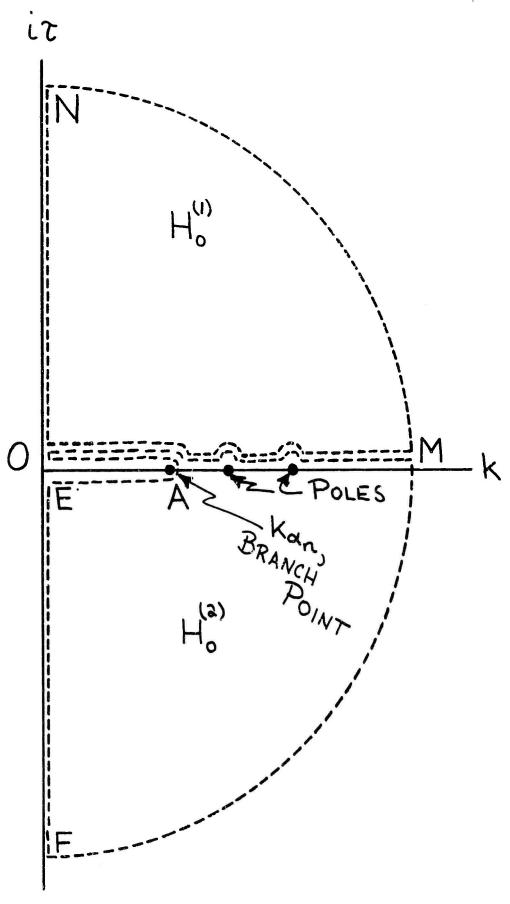


Fig. 2

Evaluation of the Integral Solutions.

The integral solutions have singularities corresponding to zeros of $\frac{A_{11}}{i\sqrt{k_{1}^{2}-k^{2}}} + \frac{A_{21}}{i\omega\rho\eta}$

These are simple poles for complex wave number and, therefore, the integral solutions are conveniently evaluated by contour integration in the complex = k + i? plane. The contours are shown in Figure 2. First the transformation

$$2J_0(3r) = H_0^{(1)}(\xi r) + H_0^{(2)}(\xi r)$$
 is made.

In addition to simple poles at

for the liquid layers and $k_{\text{am}}^2 - k_{\text{a}}^2 = 0$ and $k_{\text{am}}^2 - k_{\text{a}}^2 = 0$ for each solid layer. But since the integrands are even functions of

functions of
$$\sqrt{k_{dm}^2 - k^2}$$
 and $\sqrt{k_{pm}^2 - k^2}$

branch line integrals corresponding to branch cuts made to these branch points cancel and only the integral corresponding to the branch cut for \(\sum_{k} = \sum_{k}^{2} = \sum_{k}^{2} \) must be considered.

The Reimann surface has two sheets. To satisfy the vanishing of W_0 as ξ approaches infinity, we must remain on the sheet of the Riemann surface where the real part of ξ is greater than zero. For the contours shown in Figure 2 complex poles of

$$\frac{A_{11}}{i\sqrt{k_{a_n}^2-k^{2l}}}+\frac{A_{21}}{iwPn}=0$$

have been displaced to an unused Reimann sheet and all poles are on the real k-axis. See Ewing, Jardetzky and Press (1957, pages 135 to 137)

for the proof for a two-layer liquid half space. Along the branch cut

is pure imaginary and along the real axis from $k_{\mathbf{d}_{\mathbf{n}}}$ to

infinity
$$((-(\sqrt{3}-k_{d_n}^2))$$
 is real.

In the first quadrant $Im[(k_{4n}^{3}-k_{2}^{2})]$ is positive and

in the fourth quadrant negative. The integral for \dot{w}_{o} (omitting the \dot{c} term) is now transformed to

$$\int_{0}^{M} I_{1}d\xi + \int_{0}^{T} I_{1}d\xi = 0 (17)$$

$$C(M,N) \qquad N$$

$$\int_{E}^{And} F I_{2}d\xi + \int_{0}^{T} I_{2}d\xi + \int_{0}^{T} I_{3}d\xi = 0$$

$$C(F,M) \qquad OAE$$

$$2\pi i \sum_{0}^{T} RES(I_{2}) \qquad (18)$$

where

$$T_{i} = \begin{cases} \frac{A_{12}}{c \sqrt{k_{4n}^{2} - k^{2}}} + \frac{A_{3-2}}{c \omega \rho_{n}} | A_{5,i} \rangle_{2i} - \left[\frac{A_{1i}}{c \sqrt{k_{4n}^{2} - k^{2}}} + \frac{A_{2i}}{c \omega \rho_{n}} | A_{5} \rangle_{2i} \right] \\ \frac{A_{1i}}{c \sqrt{k_{4n}^{2} - k^{2}}} + \frac{A_{2i}}{c \omega \rho_{n}} + \frac{A_{2i}}{c \omega \rho_{n}} \\ H_{o}^{(i)} (k_{7}) \delta \end{cases}$$

$$I_{3} = \begin{cases} \frac{A_{12}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{22}}{c \omega \rho_{n}} (A_{51})_{31} - \left[\frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} (A_{51})_{32} \right] \\ \frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} (A_{51})_{32} \\ \frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} (A_{51})_{32} \\ \frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} (A_{51})_{32} \\ \frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} (A_{51})_{32} \\ \frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} (A_{51})_{32} \\ \frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} (A_{51})_{32} \\ \frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} (A_{51})_{32} \\ \frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} (A_{51})_{32} \\ \frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} (A_{51})_{32} \\ \frac{A_{11}}{c \sqrt{k_{u_{n}}^{2} - k^{2}}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} \\ \frac{A_{11}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega \rho_{n}} + \frac{A_{21}}{c \omega$$

Since $H_0^{(1)}(fr)$ approaches zero as f approaches infinity in the first quadrant, (17) is

Likewise, since
$$H_o(r)$$
 approaches zero as f approaches

infinity in the fourth quadrant (18) is

$$\int_{0}^{-i\infty} I_{a}d\xi + \int_{\infty}^{0} I_{a}d\xi + \int_{\infty}^{0} I_{a}d\xi = 2\pi i \sum_{\infty}^{0} RES(I_{a})$$
OAE

$$\int_{0}^{\infty} I_{a}d\xi = \int_{0}^{-\infty} I_{a}d\xi + \int_{0}^{\infty} I_{a}d\xi = \int_{0}^{27} AOE$$

$$-2\pi i \sum_{k=0}^{\infty} RES(I_{a}). \qquad (20)$$

Using the identity $H_0^{(1)}(\zeta r) = -H_0^{(2)}(19)$ may be written

and (20) may be written as

$$\int_{a}^{\infty} dk = -\int_{b}^{\infty} F(-cVk_{u_{n}}^{2} + e^{2t}) H_{o}^{(a)}(-cer) 2 dz$$

$$+ \int_{k_{u_{n}}}^{k_{u_{n}}} F(cVk_{u_{n}}^{2} - k^{2}) - F(-cVk_{u_{n}}^{2} - k^{2}) H_{o}^{(a)}(kr) k dk$$

$$- 2\pi i \sum_{k} RES(I_{2}).$$
Hence
$$\dot{W}_{b} = \int_{b}^{\infty} I_{i} dk + \int_{a}^{\infty} dk = \int_{b}^{\infty} (-cer) 2 dz$$

$$+ \int_{b}^{k_{u_{n}}} F(cVk_{u_{n}}^{2} + e^{2t}) - F(-cVk_{u_{n}}^{2} + e^{2t}) H_{o}^{(a)}(-cer) 2 dz$$

$$+ \int_{b}^{k_{u_{n}}} F(cVk_{u_{n}}^{2} - k^{2t}) - F(-cVk_{u_{n}}^{2} - k^{2t}) H_{o}^{(a)}(kr) k dk$$

$$- 2\pi i \sum_{k} RES(I_{2}).$$

In the first integral

$$F(i\sqrt{k_{4n}^{2}+2^{2i}}) - F(-i\sqrt{k_{4n}^{2}+2^{2i}}) = 28.$$

$$2(A_{11}A_{22} - A_{12}A_{21})(A_{51})_{21}$$

$$W\rho_{11}\sqrt{k_{4n}^{2}+2^{2i}}\left(\frac{A_{11}^{2}}{k_{4n}^{2}+2^{2i}} - \frac{A_{21}^{2i}}{w^{2}\rho_{12}^{2i}}\right)$$
but A_{11} $A_{22} - A_{12}$ $A_{21} = 1$
and therefore

and therefore
$$F(i\sqrt{k_{n}^{2}+2^{2}})-F(-i\sqrt{k_{n}^{2}+2^{2}}) = \frac{2\omega\rho_{n}\sqrt{k_{n}^{2}+2^{2}}(A_{s_{i}})_{2i}}{\omega^{2}\rho_{n}^{2}A_{i}^{2}-(k_{n}^{2}+2^{2})A_{2i}^{2}}$$

Likewise, in the second integral

Therefore, the integral solution reduces to

$$\dot{w}_{0} = \left(\frac{\omega_{1}^{2} - (\kappa_{1}^{2} + \epsilon_{2})(A_{s_{1}})_{2} H_{0}^{(2)}(-i\epsilon_{1})ede}{\omega_{2}^{2} \rho_{1}^{2} A_{1}^{2} - (\kappa_{1}^{2} + \epsilon_{2})A_{2}^{2}}\right) + \left(\frac{\omega_{1}^{2} - \kappa_{2}(A_{s_{1}})_{2} H_{0}^{(2)}(-i\epsilon_{1})ede}{\omega_{2}^{2} \rho_{1}^{2} A_{1}^{2} - (\kappa_{1}^{2} - \kappa_{2})A_{2}^{2}}\right) + \left(\frac{\omega_{1}^{2} - \kappa_{2}(A_{s_{1}})_{2} H_{0}^{(2)}(-i\epsilon_{1})ede}{\omega_{2}^{2} \rho_{1}^{2} A_{1}^{2} - (\kappa_{1}^{2} - \kappa_{2})A_{2}^{2}}\right)$$

$$-2\pi i \sum Res(I_2)$$

From the formulas
$$H_0^{(a)}(-icr) = \frac{2i}{\pi}K_0(2r)$$
 and $H_0^{(a)}(kr) = J_0(kr) - iY_0(kr)$ and multiplying $W_0^{(a)}$ by $\frac{P_0}{iwP_0}$ for a constant pressure

source of pressure amplitude Po at unit distance from the source

For the normal mode contribution,
$$-2\pi i \sum RES(\frac{P_0I_a}{iwp_s})$$
, it is

convenient to rewrite the integral for
$$w_0$$
 in the form

$$\dot{w}_0 = \left(\frac{CA_{13}}{CA_{13}} + \frac{A_{23}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right)$$

$$= \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right)$$

$$= \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right)$$

$$= \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right)$$

$$= \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right) \left(\frac{CA_{11}}{CWP_1} + \frac{A_{21}}{CWP_1} \right)$$

where the layer matrices Q are defined in the Appendix.

There are, as before, simple poles of $I_{\mathfrak{I}}$ along the real axis at

$$\frac{c\overline{A_{ii}}}{ikR_{an}} + \frac{\overline{A_{ai}}}{iwp_n} = 0 \quad \text{for } w_{e_i} | l_i l_{=1,2,...}$$

This is the period equation for the laminated half space.

From the period equation we write

From the period equation we write

$$\left(\frac{i \omega \rho_{N} C}{i k \Gamma_{\alpha_{N}}}\right)_{\omega = \omega_{0}} = \left(\frac{-A_{31}}{A_{11}}\right)_{\omega = \omega_{0}}$$
and hence
$$-2\pi i \sum_{k=k_{0}} RES \left(\frac{P_{0} I_{3}}{i \omega \rho_{s}}\right) = \left$$

But again \overline{A}_{11} \overline{A}_{22} - \overline{A}_{21} \overline{A}_{12} = 1, and therefore

$$-2\pi i \sum_{n} RES\left(\frac{P_0 I_2}{i \omega P_s}\right) =$$

$$2\pi \sum_{n} \frac{P_0 \omega (\overline{A}_{s,i})_{2i} [J_0(kr) - i Y_0(kr)]}{c^3 P_s P_n A_{ii} \frac{\partial}{\partial c} \left[\frac{c^2 \overline{A}_{ii}}{r_{an}} + \frac{\overline{A}_{3i}}{P_n}\right]} =$$

$$\frac{1}{c^3 P_s P_n A_{ii} \frac{\partial}{\partial c} \left[\frac{c^2 \overline{A}_{ii}}{r_{an}} + \frac{\overline{A}_{3i}}{P_n}\right]}{c^2 c^2 c^2}$$

In summary, the expressions for vertical particle velocity at the surface, vertical particle displacement at the surface, and pressure in a liquid layer are:

$$\dot{w}_{o} = \frac{\left(\frac{4P_{o}P_{n}Vk_{u_{n}}^{2}+2^{2}}{m\rho_{s}\left[\omega^{2}p_{n}^{2}A_{ii}^{2}-(k_{u_{n}}^{2}+2^{2})A_{2i}^{2}\right]}}{\left(\frac{2P_{o}P_{n}Vk_{u_{n}}^{2}-k^{2}(A_{s_{i}})_{2i}\left[J_{o}(kr)-iY_{o}(kr)\right]kdk}{i\rho_{s}\left[\omega^{2}p_{n}^{2}A_{ii}^{2}-(k_{u_{n}}^{2}-k^{2})A_{2i}^{2}\right]}}\right)} + 2\pi \frac{\left(\frac{P_{o}\omega(\bar{A}_{s_{i}})_{2i}\left[J_{o}(kr)-iY_{o}(kr)\right]kdk}{iP_{s}\left[\omega^{2}p_{n}^{2}A_{ii}^{2}-(k_{u_{n}}^{2}-k^{2})A_{2i}^{2}\right]}\right)}{c^{3}P_{s}P_{n}\bar{A}_{ii}\frac{\partial}{\partial c}\left[\frac{c^{2}\bar{A}_{ii}}{r_{u_{n}}}+\frac{\bar{A}_{2i}}{P_{n}}\right]}{e^{-\omega \ell}\ell_{s}}$$

$$w_{0} = \frac{w_{0}}{i\omega} = \left(\frac{4P_{0}P_{n}\sqrt{k_{a,n}^{2}+2^{2}}(A_{s,1})}{i\pi P_{s}\omega[\omega^{2}P_{n}^{2}A_{a}^{2}-(k_{a,n}^{2}+2^{2})A_{a}^{2}]}\right)$$

$$+2\pi \sum_{Q} \left\{ \frac{P_{0}\omega(\bar{A}_{S_{1}})_{2}\Gamma_{1}\Gamma_{0}kn-cY_{0}kr)}{c\omega^{3}P_{5}P_{n}\bar{A}_{11}\frac{\partial}{\partial C}\left[\frac{c^{2}\bar{A}_{11}}{r_{a_{n}}}+\frac{\bar{A}_{21}}{P_{n}}\right]} \right\}_{\substack{\omega=\omega 2\\ c=c_{Q}}}$$

$$-2\pi \sum_{\alpha} \left\{ \frac{P_{0}\omega(\bar{A}_{S_{1}})_{2}(\bar{A}_{D_{1}})_{2}[J_{0}kr)-iY_{0}kr)}{c^{4}P_{S}P_{n}\bar{A}_{1}\frac{\partial}{\partial c}\left[\frac{c^{2}\bar{A}_{1}}{C_{a_{n}}}+\frac{\bar{A}_{2}}{P_{n}}\right]_{\omega=\omega_{0}}^{\omega=\omega_{0}} \right\}$$

where for the normal mode terms for $\mathcal{P}_{\mathcal{D}_i}$ the relation $\mathcal{P}_{\mathcal{D}_i} = -(\mathcal{A}_{\mathcal{D}_i})_2 (\frac{\dot{w}_0}{C})$ has been used.

For programming it is convenient to write:

and for the absolute

value of vertical particle velocity, a quantity conveniently measured in transmission experiments,

$$|\dot{w}_{0}| = \left((-W_{1} + W_{2})^{2} + (-W_{3} + W_{4} - W_{5})^{2} \right)$$
where
$$W_{1} = \left((-W_{1} + W_{2})^{2} + (-W_{3} + W_{4} - W_{5})^{2} \right)$$

$$W_{2} = \lambda \pi \left[(-W_{1} + W_{2})^{2} + (-W_{3} + W_{4} - W_{5})^{2} \right]$$

$$W_{2} = \lambda \pi \left[(-W_{1} + W_{2})^{2} + (-W_{3} + W_{4})^{2} \right] \left[(-W_{3})^{2} + (-W_{3})^{2} \right] \left[(-W_{3})^{2} \right]$$

$$W_{3} = \left((-W_{1} + W_{2})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} \right) \left[(-W_{3})^{2} \right] \left[(-W_{3})^{2} \right] \left[(-W_{3})^{2} \right]$$

$$W_{4} = \left((-W_{1} + W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} \right) \left[(-W_{3})^{2} + (-W_{3})^{2} \right]$$

$$W_{4} = \left((-W_{1} + W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} \right) \left[(-W_{3})^{2} + (-W_{3})^{2} \right]$$

$$W_{4} = \left((-W_{1} + W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} \right) \left[(-W_{3})^{2} + (-W_{3})^{2} \right]$$

$$W_{4} = \left((-W_{1} + W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} \right) \left[(-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} \right]$$

$$W_{4} = \left((-W_{1} + W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} \right) \left[(-W_{3})^{2} + (-W_{3})^{2} \right]$$

$$W_{4} = \left((-W_{1} + W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} \right]$$

$$W_{4} = \left((-W_{1} + W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} \right]$$

$$W_{4} = \left((-W_{1} + W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} + (-W_{3})^{2} \right)$$

$$W_{4} = \left((-W_{1} + W_{3})^{2} + (-W_{1} + W_{1})^{2} + (-W_{1} + W_{1})^{2} + (-W_{1} + W_{1})^{2} \right)$$

$$W_{5} = \left((-W_{1} + W_{1})^{2} + (-W_{1} + W_{1})^{2} \right)$$

$$W_{5} = \left((-W_{1} + W_{1})^{2} + (-W_{$$

$$W_{5} = 2\pi \sum \left\{ \frac{P_{0} \omega \operatorname{Im} \left[-(\overline{A}_{S_{1}})_{a_{1}} \right] Y_{0}(kr)}{c^{3} P_{0} P_{n} \overline{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^{2} \overline{A}_{11}}{Im(r_{d_{n}})} + \frac{\operatorname{Im}(-\overline{A}_{a_{1}})}{P_{n}} \right] \right\}_{c=c_{0}}^{w=w_{0}}$$

Likewise for vertical particle displacement

$$W_{0} = -W_{0} + W_{1} - W_{8} + (W_{q} - (W_{10})^{2})$$

$$W_{0} = \sqrt{\frac{2}{9}} \frac{P_{0} P_{N} |K_{n}^{2} + Z^{2}| Im [-(A_{s_{1}})_{a_{1}}] |K_{0} | (X_{0} - X_{0})^{2}}{|K_{0}|^{2} P_{0} N |K_{n}^{2} - X_{n}^{2}| Im [-(A_{s_{1}})_{a_{1}}] |K_{0} |$$

$$W_{q} = \begin{cases} \frac{1}{\sqrt{2}} \frac{1}{\sqrt{2}}$$

and for pressure

$$W_{13} = \frac{P_0 \omega Im[-(\bar{A}_{S_1})_{2_1}]Im[-(\bar{A}_{D_1})_{2_1}]Y_0kr)}{C^4 P_5 P_n \bar{A}_{11} \frac{2}{3C} \left[\frac{C^2 \bar{A}_{11}}{Im(\Gamma a_n)} + \frac{Im(-\bar{A}_{2_1})}{P_n}\right] \omega = \omega \epsilon}$$

$$\frac{W_{15}}{2\pi} \left\{ \frac{P_{0}\omega Im \left[-(\bar{A}_{5,1})_{2,1} \right] Im \left[-(\bar{A}_{D,1})_{3,1} \right] J_{0}(kr)}{c^{4} p_{5} p_{n} \bar{A}_{11} \frac{2}{3c} \left[\frac{c^{2} \bar{A}_{11}}{Im (C_{n})} + \frac{Im (-\bar{A}_{2,1})}{P_{n}} \right]_{\omega}} \right\}$$

The layer matrices for the integral along the real axis are programmed as

$$[am)_{i}$$
 $-Im[am)_{i2}$ $(am)_{i3}$ $-Im[am)_{i4}$
 $Im[am)_{2i}$ $(am)_{2i}$ $Im[am)_{2i}$ $(am)_{24}$
 $(am)_{3i}$ $-Im[am)_{33}$ $(am)_{34}$
 $Im[am)_{4i}$ $(am)_{4i}$ $(am)_{4i}$ $Im[am)_{4i}$ $(am)_{4i}$

for solid layers and

$$[(am)_{i}] \quad [(am)_{i2}]$$

$$-Im[(am)_{2i}] \quad (am)_{22}$$

for liquid layers. For the integral along the imaginary axis the matrices are

$$\begin{bmatrix} (am)_{i_1} & (am)_{i_2} & -Im[(am)_{i_3}] & -Im[(am)_{i_4}] \\ (am)_{2_1} & (am)_{2_2} & -Im[(am)_{2_3}] & -Im[(am)_{2_4}] \\ Im[(am)_{3_1}] & Im[(am)_{3_2}) & (am)_{3_3} & (am)_{3_4} \\ Im[(am)_{4_1}] & Im[(am)_{4_2}] & (am)_{4_3} & (am)_{4_4} \end{aligned}$$

for solid layers and

for liquid layers. Likewise, for the normal mode terms the layer matrices are programmed as

for solid layers and

$$[\bar{a}_m]_{i_1} \quad \bar{I}_m[\bar{a}_m]_{i_2}]$$

$$-\bar{I}_m[\bar{a}_m]_{i_1} \quad (\bar{a}_m)_{i_2}$$

for liquid layers.

NUMERICAL COMPUTATIONS

Computer programs were written in double precision Fortran IV to compute the pressure in dynes/cm² and the vertical particle displacement in millimicrons. The programs are run on the IBM 360/91. For numerical results presented here, computing time was under six minutes. It appears that practical limits of numerical precision make the present developent most useful for the Arctic at frequencies below 50 Hz in water depths to 1 km. At higher frequencies or in greater water depths the integrands may be so oscillatory that it is difficult to achieve the desired accuracy in the numerical integrations without excessive computing time.

Computations are made in two stages. The first program, an extension of Dorman's (1962) PV 7 dispersion program, computes phase- and group-velocity dispersion, the excitation function dependent only on layering of the medium, and the excitation function for the particular source and detector depths. In the first case for the m-th normal mode the excitation function is defined by

$$\frac{2\sqrt{2\pi^2}\sqrt{\omega^2}}{\rho_n c^{\frac{3}{2}}\overline{A_n}\frac{\partial}{\partial c}\left[\frac{c^2\overline{A_n}}{Im(C_{\alpha_n})} + \frac{Im(-\overline{A_{\alpha_n}})}{\rho_n}\right]}_{\omega=\omega_m}$$

$$c=c_m$$

for pressure and
$$\frac{2 \sqrt{2\pi} \sqrt{\omega}}{\rho_{n} c^{5/2} \overline{A_{11}} \frac{\partial}{\partial c} \left[\frac{c^{2} \overline{A_{11}}}{Im(C_{n})} + \frac{Im(-\overline{A_{21}})}{\rho_{n}} \right]}$$

$$C = C_{n}$$

for vertical particle velocity. In the second case the excitation

function is
$$\left(\begin{array}{c}
2 \overline{m} \overline{w} \overline{I} \overline{m} \left[-(\overline{A}_{S_{1}})_{2,1}\right] \overline{I} m \left[-(\overline{A}_{D_{1}})_{2,1}\right] \\
P_{n} P_{S} C^{\overline{A}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
for pressure and
$$\left(\begin{array}{c}
2 \overline{m} \overline{w} \overline{u} \overline{I} \overline{m} \left[-(\overline{A}_{S_{1}})_{2,1}\right] \\
P_{n} P_{S} C^{\overline{S}_{A}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
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\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (-\overline{A}_{2,1})\right] \\
\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (\overline{A}_{n})\right] \\
\overline{P_{n} P_{S} C^{\overline{S}_{A}}} \overline{A}_{11} \frac{\partial}{\partial C} \left[\frac{C^{2} \overline{A}_{11}}{\overline{I} m (\overline{V}_{A_{n}})} + \overline{I} m (\overline{A}_{n})\right]$$$$

for vertical particle velocity. These definitions were chosen to be useful also at long ranges where $H_{n}^{(2)}(kr) =$

The second program computes the

three integrals, the normal mode contribution to the sound field, and writes and plots the absolute value of b, or w_0 as a

TABLE 1.

Model Parameters for Solensen's (1959) Computations: Source Pressure Amplitude re im,.3048 dyne/cm²: Source Depth, 15.2 m; Hydrophone Depth, 15.2m; Range, 3.048 m

Layer Thickness, m	Compressional Velocity, m/sec	Shear Velocity, m/sec	Density gm/cm ³
18.3	1463.0	0	1.03
Infinite	1609.3	0	1.24

TABLE 2.

Integral Real Axis, Real Part

Frequency, Hz	Sorensen's (1959) Value, dynes/cm ²	Our Value, dynes/cm ²		
10	.004713	.004685		
20	.030170	.030153		
40	.040278	.040485		
80	.023193	.023121		
160	021839	022201		
320	074395	074011		

TABLE 3.

Integral Real Axis, Imaginary Part

Sorensen's Value, dynes/cm ²	Our Value, dynes/cm ²		
002567	002551		
022260	022240		
061960	062274		
050986	050641		
047979	047750		
.006299	.008316		
	002567 022260 061960 050986 047979		

TABLE 4.

Integral Imaginary Axis

Frequency,	Sorensen's Value, dynes/cm ²	Our Value, dynes/cm ²
10	.095080	.095146
20	.081938	.082076
40	.041887	.041759
80	.030795	.030937
160	.030214	.030356
320	.010593	.011043

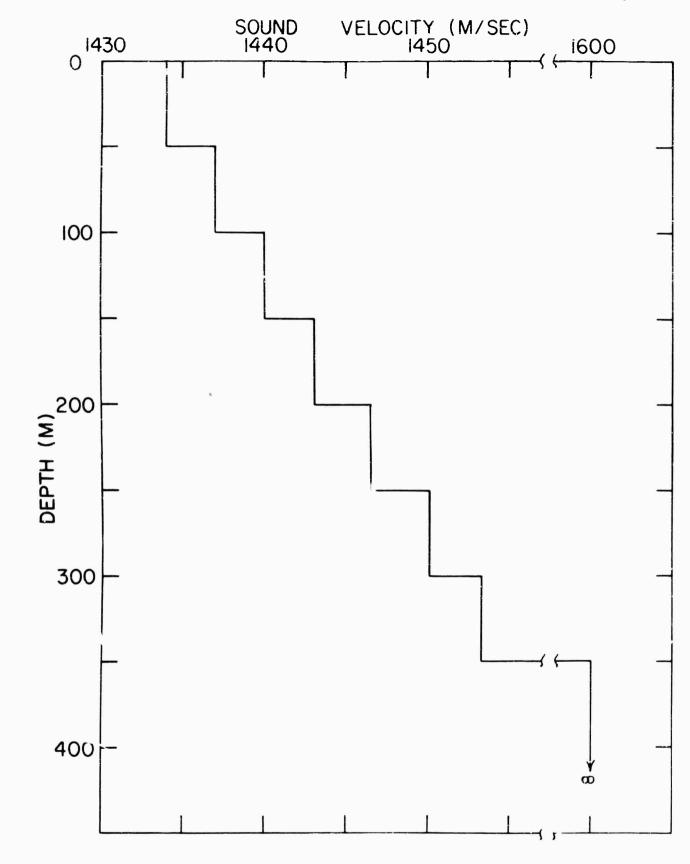
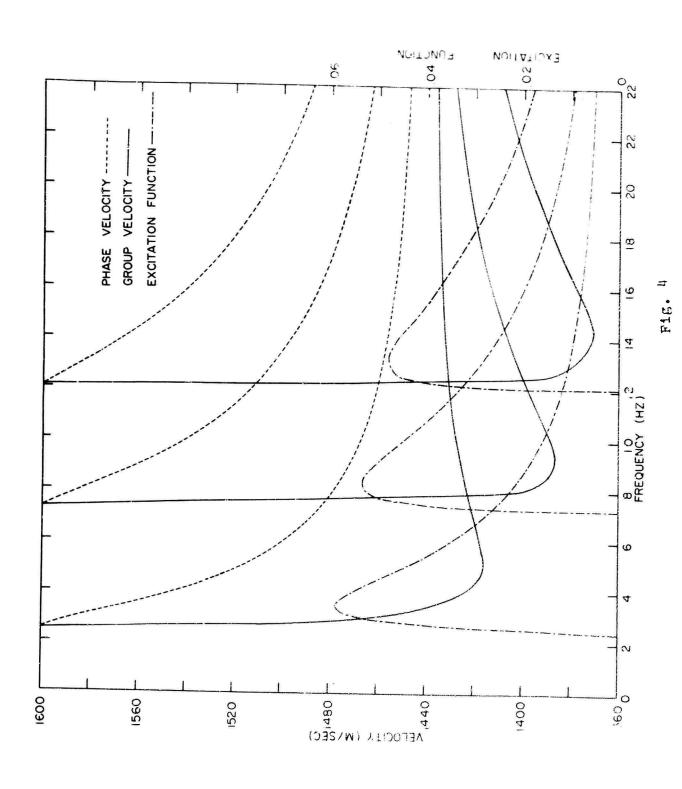


Fig. 3

TABLE 5.
Model A

Layer Thickness, m	Compressional Velocity, m/sec	Shear Velocity, m/sec	Density, gm/cm3
50	1434.0	0	1.03
50	1437.0	0	1.03
50	1440.0	0	1.03
50	1443.0	0	1.03
50	1446.0	0	1.03
50	1450.0	0	1.03
50	1453.2	0	1.03
Infinite	1600.0	0	1.20



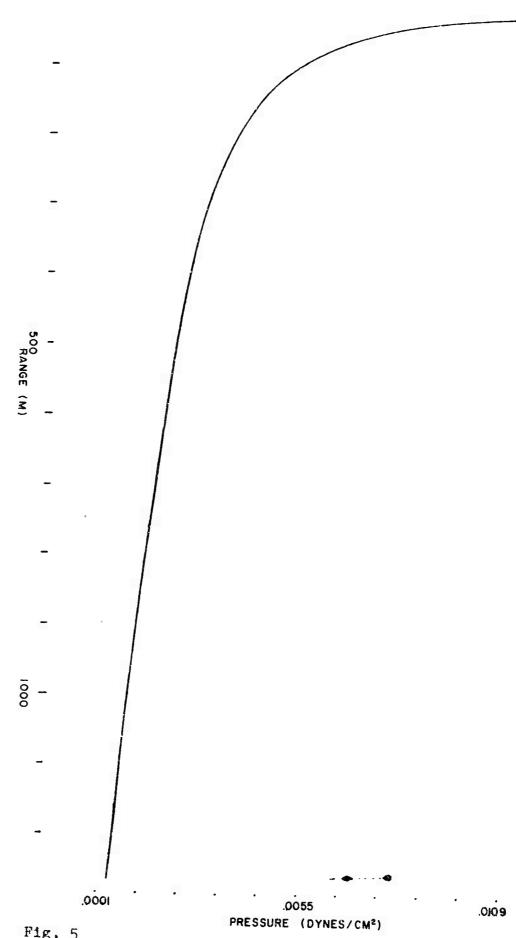
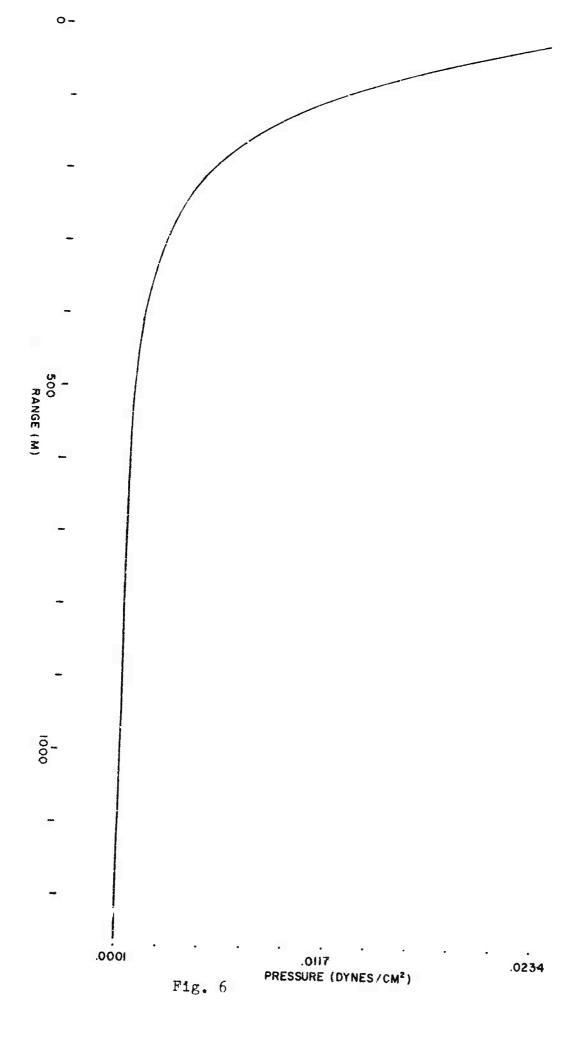
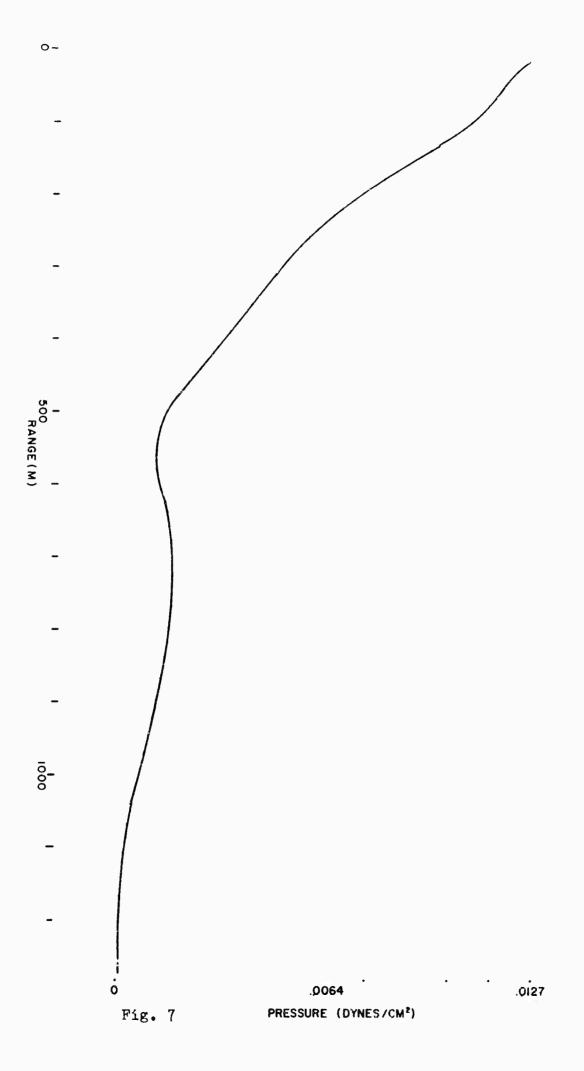


Fig. 5







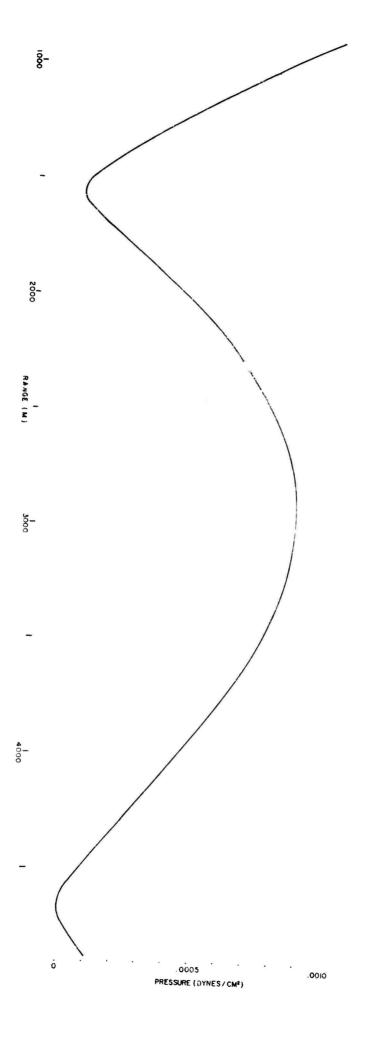


Fig. 8

Fig. 9



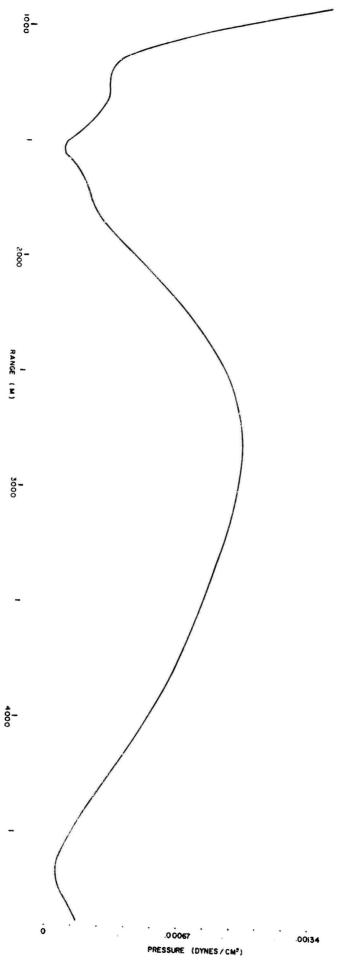


Fig. 10

function of range. In this program provision is also made to compute the three integrals as a function of frequency or detector depth at a specified range. The integrals are evaluated by Gaussian quadrature formulas (see, for example, Hildebrand, 1956). Care must be taken to employ a sufficient number of Gaussian points to obtain the desired accuracy of the integrals. A thirty-two point Gaussian quadrature formula was found to give sufficient accuracy for the computations presented here out to a range of 1500 m, but at longer ranges, higher frequencies, or greater water depths a ninety-six point formula is used. Abscissas and weight factors for the Gaussian integration are given by Davis and Rabinowitz (1956, 1958). The integration along the imaginary axis extends to infinity, but in practice a finite upper limit is chosen which gives sufficient accuracy. This may be done because the integrand is an exponential function of -kr for large k or r. The programs were checked by computations for the same two-layer model of Table 1 used by Sorensen (1959). Tables 2, 3 and 4 show that the two sets of numerical results are in very close agreement.

As an illustration of the progrms for the Arctic, computations were made for the layered model A of Figure 3, which closely follows the observed variation of sound velocity with depth. Additional parameters for the model are given in Table 5. This is a preliminary model and the ice sheet, which averages about 3-m in thickness in the central Arctic Ocean, was not included. The effects of the ice sheet on propagation are currently under investigation. Figure 4 shows phase - and group-velocity dispersion and the excitation function for pressure. At a frequency of 10 Hz only two normal modes are excited. The range dependence of the absolute

value of pressure for the normal mode contribution, the integral contribution, and the total pressure are shown in Figures 5 through 10.

ACKNOWLEDGMENTS

Dr. J. Dorman of the Lamont-Doherty Geological Observatory kindly supplied a copy of his PV-7 dispersion program. Computing facilities were provided by the Columbia inversity Computing Center, This work was supported by the U. S. Naval Ordnance Laboratory and the Office of Naval Research under contract N00014-67-A-0108-0016.

REFERENCES

- Davis, P. and P. Rabinowitz, Abscissas and weights for Gaussian quadratures of high order, J. Res. NBS, 56, 35-37, 1956.
- Davis, P. and P. Rabinowitz, Additional abscissas and weights for Gaussian quadratures of high order: Values for n=64, 80, and 96, J. Res. NBS, 60, 613-614, 1958.
- Dorman, J., Period equation for waves of Rayleigh type on a layered, liquid-solid half-space, <u>Bull Seismol Soc. Am.</u>, 52, 389-397, 1962.
- Ewing, W. M., W. S. Jardetzky and F. Press, Elastic Waves in Layered Media, McGraw-Hill, New York, 1957.
- Greene, C. R., Arctic operation of seismic transducers, Sea

 Operations Department, AC Electronics Defense Research

 Laboratories of General Motors Corporation, AC-DRL TR

 68-53, Santa Barbara, California, 1968.
- Harkrider, D. G., Surface waves in multilayered elastic media

 I. Rayleigh and Love waves from buried sources in a multilayered elastic half-space, <u>Bull. Seismol. Soc. Am., 54</u>,
 627-679, 1964.
- Haskell, N. A., The dispersion of surface waves on multilayered media, <u>Bull. Seismol. Soc. Am.</u>, 43, 17-34, 1953.
- Hildebrand, F. B., Introduction to Numerical Analysis, McGraw-Hill, New York, 1956.
- Kutschale, H., Arctic hydroacoustics, Arctic, 22, 246-264, 1969.
- Leslie, C. B., and N. R. Sorensen, Integral solution of the shallow water sound field, <u>J. Acoust. Soc. Am.</u>, <u>33</u>, 323-329, 1961.

- Fekeris, C. L., Theory of propagation of explosive sound in shallow water in Propagation of sound in the ocean, Geol. Soc. Am. Memior, 27, 1-117, 1948.
- Sorensen, N. R., Integral solution of the shallow water sound field, U. S. Naval Ordnance Laboratory, NAVORD Report 6656, White Oak, Maryland, 1959.
- Thomson, W. T., Transmission of elastic waves through a stratified solid medium, <u>J. Appl. Phys.</u>, <u>21</u>, 89-93, 1950.

Matrix elements for normal modes (see Haskell, 1953, and Dorman, 1962).

Solid layers

$$(\bar{a}_{m})_{11} = (\bar{a}_{m})_{44} = \gamma_{m} \cos \bar{P}_{m} - (\gamma_{m-1}) \cos \bar{Q}_{m}$$
 $(\bar{a}_{m})_{12} = (\bar{a}_{m})_{34} = i \left[(\gamma_{m} (\gamma_{m-1})) \sin \bar{P}_{m} + \gamma_{m} \gamma_{m} \sin \bar{Q}_{m} \right]$

$$(C_{m})_{13} = (\bar{Q}_{m})_{24} = -(p_{m}c^{2})^{T}(cosP_{m} - cos\bar{Q}_{m})$$
 $(\bar{Q}_{m})_{14} = c'(p_{m}c^{2})^{T}[\Gamma_{a_{m}}^{-1} sinP_{m} + \Gamma_{B_{m}} sin\bar{Q}_{m}]$
 $(\bar{Q}_{m})_{21} = (\bar{Q}_{m})_{43} = -c[\Gamma_{d_{m}} r_{m} sin\bar{P}_{m} + \Gamma_{B_{m}} r_{m} sin\bar{P}_{m} r_{m} r_{m} r_{m} sin\bar{P}_{m} r_{m} r_{m} r_{m} sin\bar{P}_{m} r_{m} r_{m}$

$$(\bar{q}_{m})_{33} = (\bar{q}_{m})_{33} = -(\bar{r}_{m} - 1)\cos\bar{p}_{m} + \bar{r}_{m}\cos\bar{q}_{m}$$

$$(\bar{q}_m)_{ij} = (q_m)_{44} = \cos P_m$$

$$(\bar{q}_m)_{ij} = \frac{i r_{dm}}{P_m c^2} \sin P_m$$

$$(\bar{q}_m)_{aj} = \frac{i p_m c^2}{r_{dm}} \sin P_m$$
Matrix elements for integrals elements

Matrix elements for integrals along real axis:

$$(a_m)_{ij} = (a_m)_{44} = 2(\frac{k}{k_{\beta m}})^2 \cos P_m - (2(\frac{k}{k_{\beta m}})^{-1}).$$

$$(a_m)_{12} = (a_m)_{34} = i \left[\frac{k(2(k_{pm})^2 - 1)}{\sqrt{k_{pm}}} \frac{sinP_m}{\sqrt{k_{pm}}} + \frac{1}{\sqrt{k_{pm}}} \frac{sinP_m}{\sqrt{k_{pm}}} \right]$$

$$(Q_m)_{22} = (Q_m)_{33} = \left[-\left(2(\frac{k}{k_{Bm}})^2 - 1 \right) \cos P_m + 2(\frac{k}{k_{Bm}})^2 \cos Q_m \right]$$

$$[COSP_m - COSQ_m]$$

$$(Q_m)_{32} = \frac{ip_m \omega}{k} \left[\left(2\left(\frac{k}{kp_m}\right)^2 - 1\right)^2 \frac{KSINP_m}{\sqrt{k_m^2 - k^2}} + \frac{1}{k^2 - k^2} \right]$$

$$2(\frac{k}{k}g_{m})^{2}(\frac{k}{k}g_{m}-k^{2})^{2}(\frac{k}{k}g$$

Solid layers
$$P_{m} = h_{m} \sqrt{k_{m}^{2} + 2^{2}}$$

$$Q_{m} = h_{m} \sqrt{k_{m}^{2} + 2^{2}$$

$$(a_m)_{21} = (a_m)_{43} = \left[\frac{2c}{k_{Bm}}\sqrt{k_{A_m}^2 + 2^{27}} SINP_m - \frac{2c}{k_{Bm}}\sqrt{k_{A_m}^2 + 2^{27}}\right]$$

$$(Q_m)_{22} = (Q_m)_{33} = (2(\frac{2}{k_{\text{Rm}}})^2 + 1) \cos P_m - 2(\frac{2}{k_{\text{Rm}}})^2 \cos Q_m$$

$$(2m)_{32} = i \omega p_m \left(2 \left(\frac{2}{k_{pm}} \right)^2 + 1 \right)^2 \frac{SINP_m}{\sqrt{k_{pm}^2 + 2^2}} - \frac{64}{\sqrt{k_{pm}^2 + 2^2}} \left(\frac{2}{k_{pm}} \right)^2 \left$$

Liquid layers

$$P_{m} = h_{m} V_{k_{m}}^{2} + z^{2l}$$

$$(Q_{m})_{11} = (Q_{m})_{44} = COSP_{m}$$

$$(Q_{m})_{12} = iV_{k_{m}}^{2} + z^{2l} SINP_{m}$$

$$(Q_{m})_{12} = iP_{m} SINP_{m}$$

$$(Q_{m})_{21} = iP_{m} SINP_{m}$$

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4 DESCRIPTIVE NOTES (Type of report and inclusive dates)			
Technical Report			
5 AUTHOR(S) (First name, middle initial, last name)			
Henry W. Kutschale			
			!
6 HEPORT DATE	76. TOTAL NO. OF	PAGES	76. NO. OF REFS
February 1970	64		13
BE. CONTRACT OR GRANT NO.	98. ORIGINATOR'S	REPORT NUM	BER(S)
N00014-67-A-0108-0016			
& PROJECT NO.		1	
NR 307-320/1-6-69 (415)			
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	this report)		
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	ILS. Naval	Ordnand	ce Laboratory,
	1		Spring, Maryland
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of layers. The integral over wave	number has	s singul	arities in the
integrand and is conveniently tran	sformed in	to the c	omplex plane. By
a proper choice of contours, compl			
sheet of the two-leaved Riemann su	-	_	
for the multilayered system reduce			
sum of two integrals, one along th			_
imaginary axis. Both integrals ar	e evaluate	d by a G	aussian quadrature
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tion in the Arctic Ocean sound cha			reliminary compu-
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tations and the ice layer, which a			
is not included in the layered sys		/	of the ice layer on
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KEY WORDS	ROLI	WT	ROLE	wT	ROLE	wt	
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Hydroacoustics							
Sound field							
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Matrix							
Integral solution							
Digital computer							
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